

THE GEOLOGY AND STRUCTURE OF NEOGENE ROCKS IN DADYAL AND ADJACENT AREAS, IN THE SUB-HIMALAYAS, AZAD JAMMU AND KASHMIR, PAKISTAN

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ABSTRACT

The project area lies along the western limb of the Hazara Kashmir Syntaxis. The Hazara Kashmir Syntaxis is an antiformal structure. The project area includes Dadyal, Bihari, Ratta, Sammorthala, Samlotha, Dhinger, Panwar and Karara Saiyidan. The Late Miocene to Recent Himalayan Neogene rocks are exposed in the core of the syntaxis. These Neogene rocks include the Kamlial Formation, Chinji Formation, Nagri Formation, Dhok Pathan Formation, Soan Formation and Recent alluvium. The major folds in the project area include the Mirpur syncline, Pheran Dinpur anticline and Sadaqabad Syncline. The folds are northwest-southeast trending and southwest or northeast vergent, asymmetric and open. The only fault in the project area is Ratta Fault. The Ratta Fault is a reverse fault. The northwest southeast trending folds and faults are the result of northeast southwest Himalayan compression in the area.

1. Introduction

The project area in Dadyal, Mirpur district of Azad Jammu and Kashmir, Pakistan is confined in the east, west and south by the Riasi thrust, Jhelum fault and Salt Range thrust respectively (Fig. 1). The location of study area in the western limb of the Hazara Kashmir Syntaxis is shown on the tectonic map (Fig. 1).

The Dadyal, Bihari, Ratta, Sammorthala, Dhinger, Panwar, and Karara Saiyidan area are mapped on the toposheet 43/G-11, Survey of Pakistan. The areas lying at coordinates 73° 38' 50"E to 73° 45' 00" E and 33° 15' 00" N to 33° 23' 00" N are well known for fertile land, forests and the frequent occurrence of mountains. The Jhelum river in the area has been harnessed to form Mangla Reservoir.

The Nagri Formation, Dhok Pathan Formation, Soan Formation, Mirpur Formation and recent alluvium are well exposed in the area.

2. Previous work

The Mirpur and adjoining areas were previously surveyed by different geologists (Ledekker, 1876; Wadia, 1928; Chuhdary and Ashraf, 1983; Chaudhary and Ashraf, 1984; Baig and Lawrence, 1987; Wells and Gingerich, 1987; Akhtar et al., 2004; Munir and Baig, 2006) before the commencement of project in the area (Yasin, 2012).

3. Materials and methods

The city, adjacent areas of Mangla reservoir, recondite places and lonely desolate valleys were surveyed in mild weather with utmost care during 30 days field work. The field photographs (Fig. 5) were taken by means of a digital camera. The rock Formations were mapped by using Brunton compass during onerous task (Fig. 3). In addition, the instrument was also used to record the attitude of bedding (Fig. 3). The route and traverse, geological and structural maps of the 75km² area on 1:50,000 scale were compiled (Fig. 2). Moreover, the detailed structural cross-sections and β and π diagrams were prepared for the structural analysis of the area (Fig. 5). The primary sedimentary structures remained helpful

in the identification of the facing of rock units (Fig. 4a). The data used in structural interpretation has been enumerated in the tables 1 & 2.

4. Results and Discussions

4.1 Stratigraphy

The geological succession of the Dadyal and adjacent areas of Mirpur is presented in Table 3. The rocks of the project area (Table. 4) are described as follows:

4.1.1 Siwalik Group

Meddicot, 1864 first used the term "Siwalik" to describe the upper part of the "Sub-Himalayan System". The Siwalik Group is comprised of Neogene rocks, deposited during the Himalayan orogeny. The Siwalik Group contains arenaceous, argillaceous and conglomeratic material. The following rock sequence of Siwalik group is exposed in the project area.



Figure 1. Regional tectonic map of the northwest Himalayas of Pakistan after (Baig and Lawrence, 1987; Monalisa and Khawaja, 2004). The rectangle shows the location of the project area.

- Soan Formation
- Dhok Pathan Formation
- Nagri Formation

4.1.1.1 Nagri Formation

The "Nagri Formation" was accepted by the Stratigraphic committee of Pakistan after Lewis (937). The Nagri Formation is mapped in Goda Sultana, Karara Saiyidan and Jinour areas (Fig. 2; Fig. 3). The sandstone and clays (60 to 40%) are the dominant lithologies of the Formation. The sandstone is greenish gray, weathered to light gray and brownish gray in outcrops. In addition, the sandstone is medium to coarse grained, hard, compact, medium to thick bedded and massive. Furthermore, the sandstone is cross bedded and exhibits salt and pepper texture (Fig. 4b). The clays vary in colour from cream to light gray. The clays are soft and friable. The mudstone and gritty clays are in scanty amount.

The minerals quartz, feldspar, biotite, epidote, muscovite, tourmaline, plagioclase and garnet coexist in the rock. The biotite is in excess amount. The pink garnet and Panjal volcanic clasts are highly concentrated towards the top of the Formation. The clasts of granite, quartzite and marble are also disseminated in the outcrop. The formation has a high concentration of epidote and biotite than Kamlial Formation. The formation has gradational contacts with the Dhok Pathan and Chinji Formations in the upper and lower halves respectively.

The earlier workers assigned Late Miocene age to the Formation (Shah, 2009).

4.1.1.2 Dhok Pathan Formation

The term "Dhok Pathan" of Pilgrim (Pilgrim, 1913) was redefined by Cotter (1913) as the "Dhok Pathan Formation" and finally approved by the Stratigraphic Committee of Pakistan (Shah, 2009).

The Formation is mapped in Bihari, Karara Sayidan, Khurd Piran Goda and Panwar areas (Fig. 2; Fig. 3). The sandstone, clay (50 : 50) and hard compact conglomeratic beds form the composition of rock. The sandstone is medium to coarse grained, soft, porous, friable (Fig. 4d) and less compact as compare to the sandstone of Nagri Formation. The conglomeratic levels are thick, hard and compact. The cross bedding (Fig. 4a), rip ups, lenticular bedding and ribbed topography (Fig. 4c) are the non tectonic structures of Dhok Pathan Formation. The clays are red and maroon in outcrops. The clasts vary in size from 2mm to 5cm diameter.

The minerals quartz, feldspar, muscovite, pink garnet, brown garnet, epidote, minor biotite and tourmaline form the mineral composition of the Dhok Pathan Formation. The ribbed topography (Fig. 4c) and presence of pink garnet distinguish the Dhok Pathan Formation from the rest of Formations.

The contacts with the overlying and underlying formations are gradational. However, in Ratta the formation has a faulted contact with Soan Formation (Fig 2; Fig. 3). The earlier workers declared Late Miocene age to the Formation (Shah, 2009).

4.1.1.3 Soan Formation

The Stratigraphic Committee of Pakistan has approved the name Soan Formation following Kravchenko (1964).

The sandstone, dark gray shale and conglomerates are the dominant constituents of the Soan Formation. The compact, thick, hard and massive conglomerate levels are ubiquitous near Dingar, Anb and Dadyal. The conglomerates include the pebbles and boulders of quartzite, Panjal volcanics, Murree sandstone, Paleocene-Eocene limestone, cherty dolomite and granite gneisses. The conglomerates are tessellated in arenaceous matrix (Fig. 4e). The pebbles and boulders have rounded to sub rounded outlines, few mm to 30 cm diameters and the thickness of about 0.5 meter has been paced out in the lenticular sandy layers (Fig. 4e). The sandstone lenses are exposed in Siakh (Fig. 4e), near Samlotha, Pagal Chak and Dadyal.

The contact of Soan Formation with Dhok Pathan Formation is faulted in project area (Fig. 4h), gradational in other areas and sharp (angular) with Mirpur Formation in the Mirpur area. The earlier workers proposed Pliocene age to the Formation (Shah, 2009).

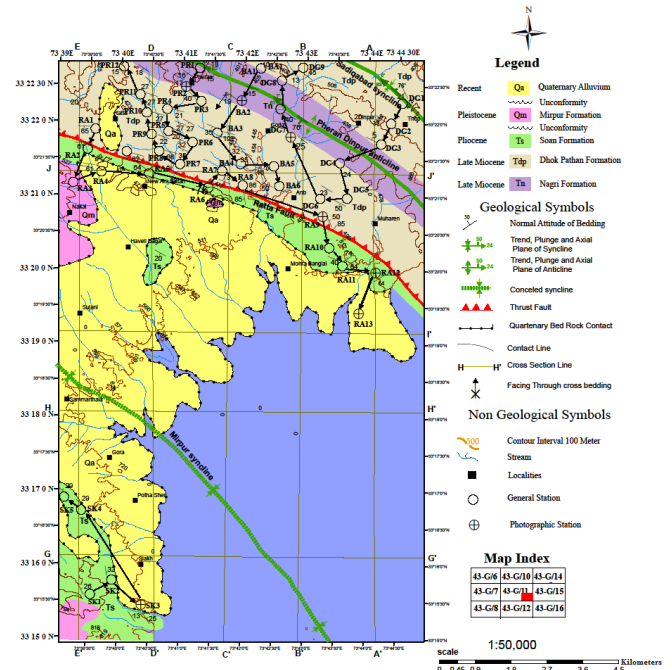


Figure 2. Route and Traverse Map of Ratta, Behari, Dadyal area, District Mirpur, Azad Kashmir, Pakistan.

4.1.1.4 Mirpur Formation

The Formation is mapped in the Sammorthala, Ratta, Siakh, Nakai and Dadyal areas (Fig. 2; Fig.3). The formation is a Pleistocene flood or channels fill deposit. The framework grains are loosely packed within the argillaceous matrix (Fig. 4f). The clast varies in size from 2mm to 30cm diameter (Fig. 4f).

These conglomerates followed the topography of the underlying formation and appear as a blanket over the bed rocks. The upper contact with recent alluvium and lower contact with Soan Formation are unconformable. The Mirpur Formation can be distinguished from the Soan formation by argillaceous cement. The age of the formation is Pleistocene (Shah, 2009).

4.1.1.5 Recent Alluvium

The fresh water deposits are mapped in the Mohra Bangial, Ratta, Sammorthala, Siakh, Potha Bangash, Potha Sher, Mughal Pura and Bihari areas (Fig. 2; Fig. 3). The deposits contain rounded to subrounded gravel, sand, silt and clay. These deposits are less compact (Fig. 4g). The alluvium marks an unconformity on the top of Mirpur Formation.

4.2 Structure

The project area is surrounded by the Jhelum Fault, Riasi Fault and Salt Range Thrust in the western, eastern and southern directions, respectively (Fig. 1). The area is exposed in the south of Hazara Kashmir Syntaxis and is highly deformed by Himalayan orogeny (Fig. 1). The Hazara Kashmir Syntaxis (Fig. 1) is an antiformal structure formed by the abrupt change in northwest regional strike of the Himalayan thrust sheets to the northeast. The Main Boundary Thrust and Panjal Thrust show contortion around the Hazara Kashmir Syntaxis (Fig. 1). The folds in the project area are northwest-southeast trending, open and northeast or southwest vergent (Tables, 1 & 2).

4.2.1 Structure of the area

The Tertiary to Recent cover sequence of Indian Plate is deformed into folds and faults (Fig. 1). The structures is described as follows:

4.2.2 Folds

There are three following major folds in the area.

- Pheran Dinpur Anticline
- Sadiqabad Syncline
- Mirpur Syncline

4.2.2.1 Pheran Dinpur Anticline

The Pheran Dinpur anticline is formed by the folding of Nagri and Dhok Pathan Formations (Fig. 2; Fig. 3). The Nagri Formation forms the core whereas the Dhok Pathan Formation forms the limbs of the fold (Fig. 5, cross section JJ'). The southwestern limb of the Pheran Dinpur anticline is displaced by the Ratta fault (Fig. 5, cross sections AA', BB', CC', DD', JJ').

The Chinji and Kamlial Formations confront in the subsurface (Fig. 5, cross-sections AA', BB', CC', DD', JJ'). The strike of northeast limb displays orientation in N32°W to N79°W whereas the strike of the southwestern limb exhibits orientation in N29°W to N58°W direction. The northeastern limb of fold dips at 10°NE to 50°NE whereas the southwestern limb of fold dips at 24°SW to 40°SW (Table. 1). The attitude of the axial plane is slanted in N30°W 80°SW to N72°W/77°SW directions (Table. 1; Fig. 6.).

The fold axis of the anticline trend and plunge at 5°/330° to 7°/288° (Table, 1). The interlimb angle of the fold varies between 106° to 120°. On the basis of interlimb angle, the fold is classified as an open fold. The Pheran Dinpur anticline is northeast verging fold. The regional anticline is interpreted as hanging wall anticline, formed along the Ratta Fault (Fig. 5; cross sections AA', BB', CC', DD', JJ').

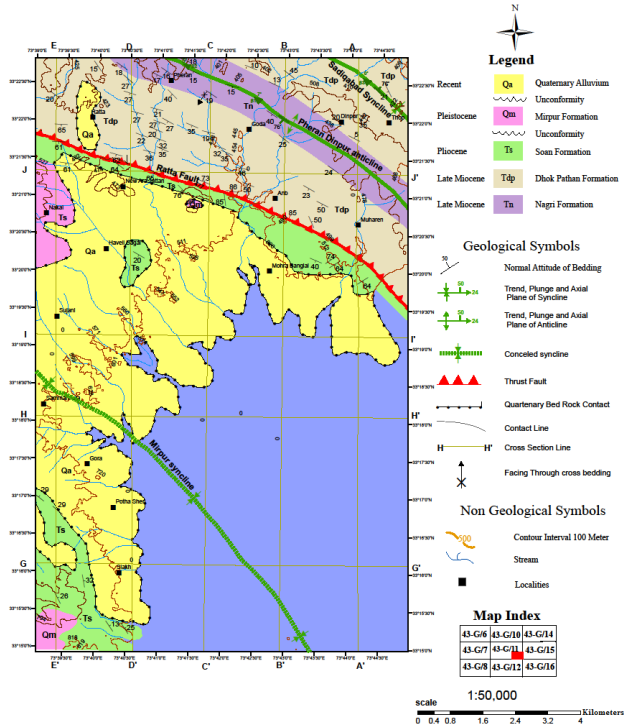


Figure 3. The composite Geological and structural map of Ratta, Behari and Dadyal areas, District Mirpur, Azad Kashmir, Pakistan.

4.2.2.2 Sadiqabad Syncline

The Sadiqabad syncline is a regional syncline formed by the folding of Dhok Pathan and Nagri Formations. The Dhok Pathan Formation lies in the core whereas Nagri Formation forms the limbs (Fig. 5, cross sections AA', BB', CC', DD')

The strike of the northeastern limb is oriented in N26°W whereas the strike of the south western limb is oriented in N30°W direction. The northeastern limb of syncline dips at 08°SW whereas southwestern limb dips at 15°NE. The Sadiqabad syncline exhibits a trend and plunge of 0°/334°. The attitude of axial plane falls in N26°W /85°SW direction Table. 2; Fig. 6). The Sadiqabad syncline has an interlimb angle of about 157° (Table, 2; Fig. 6). On the basis of interlimb angle the fold is classified as gentle fold. The Sadiqabad syncline is northeast vergent fold.

4.2.2.3 Mirpur Syncline

Mirpur syncline is regional syncline formed by the folding of Soan Formation. In project area, the Mangla reservoir and Quaternary alluvium has concealed the syncline. (Fig. 2; Fig. 3: cross sections AA', BB', CC', DD', JJ').

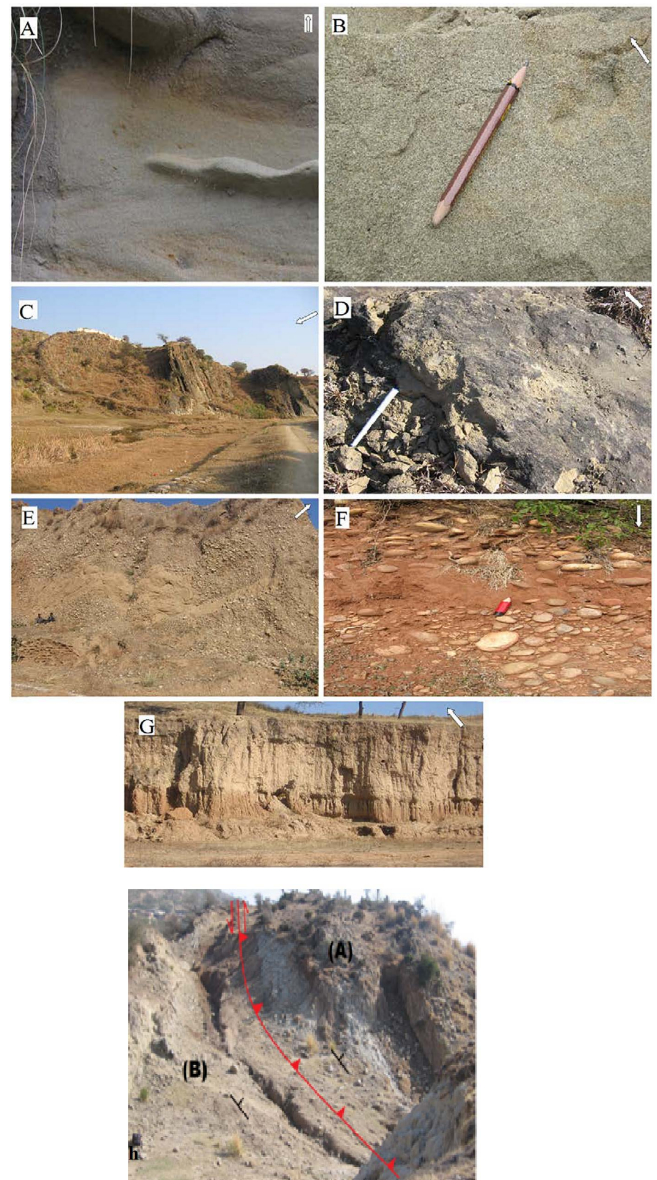
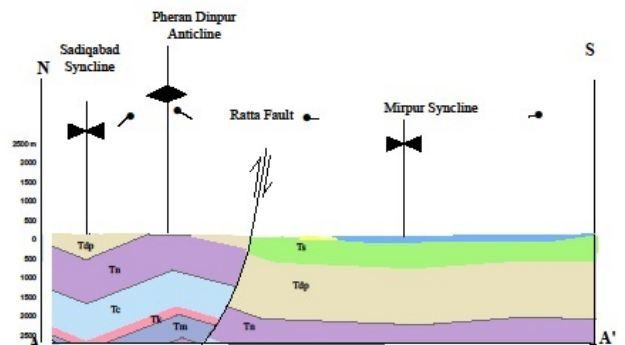


Figure 4. (a) Photograph showing cross bedded sandstone of Nagri Formation near Pheran (photo facing south), (b) showing salt and pepper texture in the sandstone of Nagri Formation near Pheran (Photo facing southwest), (c) Photograph showing ribbed topography (cyclic bedding) in Dhok Pathan formation near Goda (Photo facing southwest), (d) Photograph showing loosely compacted sandstone of Lower Soan Formation, Mohra Bangial (Photograph facing southwest), (e) Photograph showing lenses of tessellated conglomerate in upper Soan Formation near Siakh (Photo facing south east), (f) Photograph showing conglomerates of Mirpur Formation near Nakai (Photo facing North), (g) Photograph showing thick layer of recent alluvium near Mohra Bangial (Photo facing southwest), Photograph showing Ratta fault between Dhok Pathan Formation (A) and Soan Formation (B) near Anb (Photograph facing south).



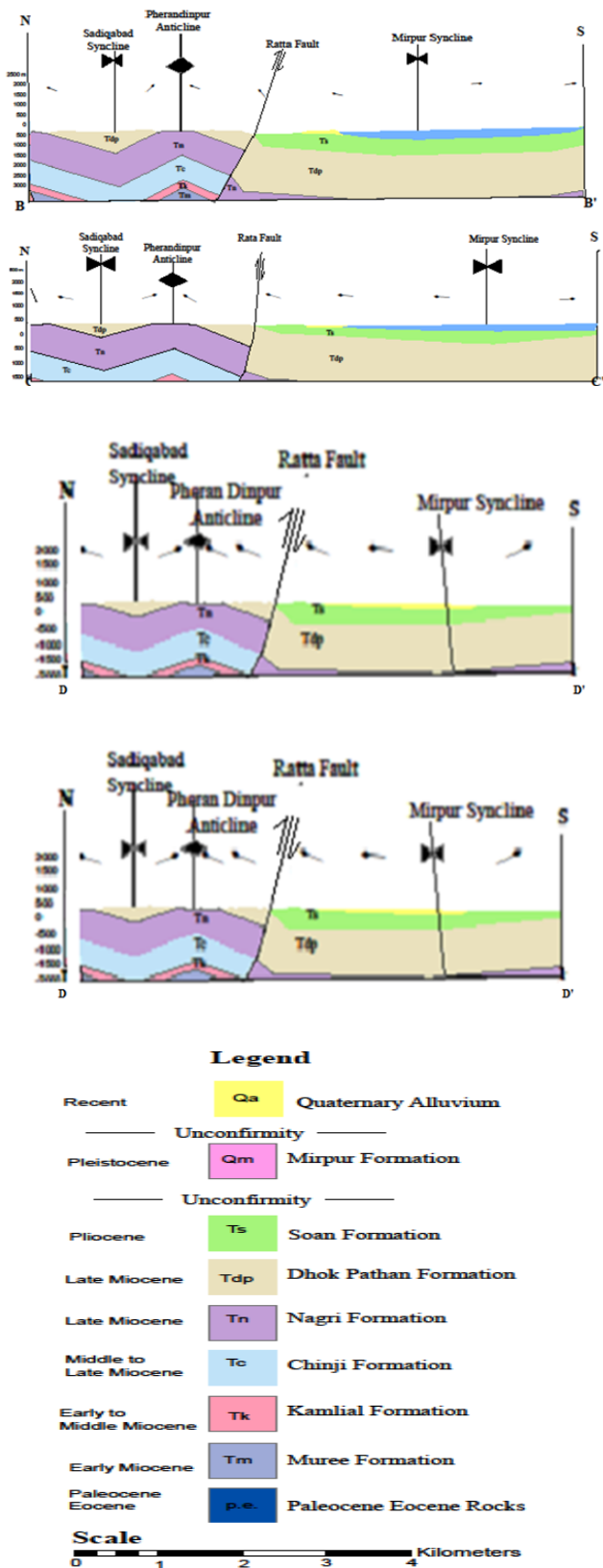


Figure 5. AA', BB', CC', DD', JJ', Structural cross sections of Sadiqabad, Ratta area and adjacent areas through the profiles on map, District Mirpur, Azad Kashmir, Pakistan.

4.2.2.4 Faults

There is only one major fault in the area called as Ratta Fault.

4.2.2.5 Ratta Fault

The Ratta Fault is a northwest-southeast trending reverse fault and has displaced the southeastern limb of Pheran Dinpur anticline (Fig. 2; Fig. 3). In the project area, the Dhok Pathan Formation thrusts over the Soan Formation (Fig. 2; Fig. 3; cross sections AA', BB', CC', DD' and JJ'). The attitude of the hanging wall falls in N45°W/23°NE to N55°W/35° NE directions and the attitude of footwall block falls in N48°W/32°NE

directions. The attitude of the fault plane is about N70°W/85°NE (Fig. 4h). The Ratta Fault is a reverse fault (Fig. 4h).

The shearing, crushing (Fig. 4h), fault breccia, gouge and drag folding are the clues of deformation along the Ratta Fault.

Table 1. Structural interpretations of Pheran Dinpur Anticline.

Attitude of bedding		Axial plane	Trend and Plunge of Fold axis	Interlimb Angle	Fold Type
Northeastern Limb	Southwestern Limb				
(β ₁) N79°W/50°NE	N58°W/24°SW	N72°W/77°SW	7°/288°	106°	Open
(β ₂) N32°W/10°NE	N29°W/40°SW	N30°W/80°SW	5°/330°	120°	Open

Table 2. Structural interpretations of Sadiqabad Syncline.

Attitude of Bedding		Axial Plane	Trend and Plunge of Fold axis	Interlimb Angle	Fold Type
Northeastern Limb	Southwestern Limb				
(β ₂) N26°W/08°SW	N30°W/15°NE	N26°W/85°W	0°/334°	157°	Gentle

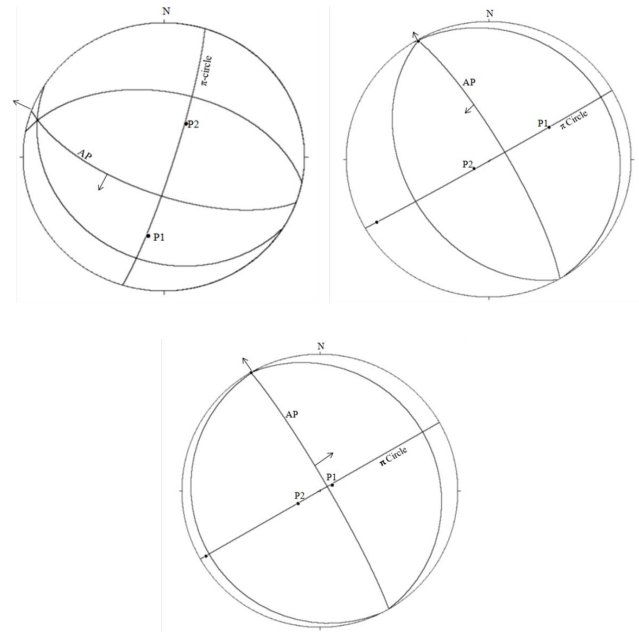


Figure 6: The π and β diagram of Pheran Dinpur anticline (1st and 2nd) and β diagram of Sadiqabad syncline (3rd).

Table 3. The Geological succession of the Mirpur and adjoining areas (Yasin, 2012).

Formation	Age	Description
Quaternary alluvium	Recent	Consist of unconsolidated deposits of clay silt, gravel.
Unconformity.....		
Mirpur Formation	Pleistocene	Poorly sorted conglomerates consisted of pebbles, cobbles of igneous, metamorphic and sedimentary rocks.
Unconformity.....		
Soan Formation	Pliocene	Conglomerates, Brown, yellow gray clays and claystone, gray sandstone.
Dhok Pathan Formation	Late Miocene	Dominantly consists of sandstone, siltstone and clays. Sandstone is grey, fine to medium grained and medium to thick bedded.
Nagri Formation	Late Miocene	Dominantly, consists of greenish grey sandstone, siltstone and mudstone. Sandstone has massive beds and has medium to coarse grained texture. Sandstone alternates with clay and are 60% and 40% respectively.
Chingi Formation	Middle to late Miocene	Red to purple, greenish grey, ash grey sandstone and siltstone and purple and reddish brown mudstone. 60% clays and 40% sandstone.
Kamilial Formation	Early to middle Miocene	Mainly sandstone, clays and intraformational conglomerates.
Murree Formation	Early Miocene	Mostly clays, shales and sandstone. Sandstone is red to purple red in colour and is fine to medium grained.
Kuldana Formation	Middle to Late Eocene	Variegated shales with subordinate sandstone. Shales are arenaceous.
Chorgali Formation	Early Eocene	Mostly Calcareous shale, nodular limestone and dolomitic limestone, clays and quartzite.
Margalla Hill Limestone	Early Eocene	Main nodular fossiliferous limestone with shales.
Patala Formation	Late Paleocene	Mainly shales with interbedded limestone.
Lockhart formation	Early Paleocene	Grey to dark grey limestone with subordinate shales.
Hangu Formation	Early Paleocene	Mainly Laterite, bauxite and fireclay.
Unconformity.....		
Muzaffarabad Formation	Cambrian	Mainly Dolomitic limestone with cherty dolomite and chert bands.
Unconformity.....		
Dogra Formation	Precambrian	Slates, phyllite and shales with limestone.

Table 4. The stratigraphic sequence of the study area.

Formation	Age	Description
Alluvium	Recent	Unconsolidated deposit of clay, gravel, pebble etc.
Unconformity.....		
Mirpur Formation	Pleistocene	Poorly sorted conglomerates consisted of pebbles, cobbles of igneous, sedimentary and metamorphic rocks.
Unconformity.....		
Soan Formation	Pliocene	Conglomerate levels are high very thick, hard and compact. Conglomerates constitute the fragments of cherty dolomite, panjal volcanics, granite, gneisses and compact arenaceous matrix. Bentonite is present in lower bed.
Dhok Pathan Formation	Late Miocene	50% sandstone and 50% clay, cyclic bedding, ribbed topography, flaser or lenticular bedding presents, hard and compact conglomerate beds are present, pink garnet also appears.
Nagri Formation	Late Miocene	70% sandstone and 30% clay present. Sandstone is hard and compact, volcanic clasts are also present in upper part. Spheroidal weathering is also present.

5. Discussion

The project area exists in the southern part of Hazara Kashmir Syntaxis. The Hazara Kashmir Syntaxis is an antiformal structure, formed by the late Tertiary Himalayan collision between the Indian and Eurasian plates. The southeastern limb of the Hazara syntaxis is imbricated along Punjal Thrust, Main Boundary Thrust and Riasi Fault (Baig and Lawrence, 1987). The western limb of the Hazara Kashmir Syntaxis is displaced by left lateral strike slip Jhelum Fault.

The core of syntaxis is comprised of a deformed zone. The deformed zone makes a domal structure due to the layer compression (Bossart et al., 1988). The Himalayan neogene rocks form the core of syntaxis. These rocks represent the part of cover sequence of Indian plate.

The Late Miocene to recent deposits in the study area include the Nagri Formation, Dhok Pathan Formation, Soan Formation, Mirpur Formation and Recent alluvium. The area is highly deformed into regional folds and faults. The folds and faults are northwest southeast trending. The folds of the area are southwest or northeast vergent. The folds are open to tight in nature. The major folds include Pheran Dinpur anticline, Sadiqabad syncline and Mirpur syncline. The Pheran Dinpur anticline is formed by the folding of Nagri Formation (Fig. 2; Fig. 3; Cross sections, AA', BB', CC' and JJ'). The southwestern limb of

the anticline is relatively steeper than the northeastern limb (Table. 1). The Pheran Dinpur anticline is an open and southwest verging asymmetric fold. The Pheran Dinpur anticline is a plunging fold.

The Sadiqabad syncline is a regional syncline formed by the folding of Dhok Pathan Formation and Nagri Formation. The Dhok Pathan Formation is in the core whereas Nagri Formation lies on the limbs (Fig. 2; section AA').

The Ratta Fault is a reverse fault in the project area. The Ratta Fault is a northwest-southeast trending fault and cuts the southeastern limb of Pheran Dinpur anticline. In the project area, the Dhok Pathan Formation is thrust over the Soan Formation (Fig. 3; cross sections, AA', BB', CC', DD' and JJ'). The attitude of the hanging wall block ranges from N45° W/23°NE to N55°W/35°NE and the attitude of footwall block is N48° W/32°NE. The attitude of the fault plane is N50°W/61°NE. The Ratta Fault is a reverse fault (Fig. 4h).

The shearing and crushing are also present along the fault zone (Fig. 4h). The fault breccia and gouge can be identified in the project area along the Ratta Fault. The drag folds are common along this fault.

1. Conclusion

The project area lies in the southeastern part of Hazara Kashmir Syntaxis. The Hazara Kashmir Syntaxis is an antiformal structure and the part of Himalayan foreland folds-and-thrust belt. The core of syntaxis possesses metamorphic and sedimentary rocks. The sedimentary rocks are mainly Neogene molasse deposits, formed by the late Tertiary Himalayan Orogeny.

In the project area, the thick cover sequence of sedimentary rocks ranging from Late Miocene to Recent age is exposed. These Neogene rocks are result of the collision of the Indian and Eurasian plates. The uplift of the Himalayan sediments caused erosion, transportation and deposition of these rocks in front of rising Himalayas. The project area is deformed due to the regional folds and faults. The fault is northwest-southeast trending. The fault is reverse in nature. The folds of the area are northwest or southeast vergent and asymmetric in nature. These folds are open. These are northwest or southeast plunging folds. The major folds of the area are Pheran-Dinpur Anticline, Sadiqabad Syncline and Mirpur Syncline.

The major fault in the area is Ratta Fault. It is a splay fault of Jehlam Fault. It is a reverse fault.

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